# Crust to Core workshop: An introduction to Perple\_X

Part 2: The structure of a Perple\_X calculation

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## Perple\_X:

- Is written and maintained by Jamie Connolly (ETH Zürich)
- Is written in FORTRAN, with the source code available.
- Took quite a long time to write...
- Accepts thermodynamic data from any source, provided it is formatted correctly (in a simple text file).

Connolly, J.A.D., Kerrick, D.M., 1987. An algorithm and computer program for calculating composition phase diagrams. *Computers and Geosciences* 11, 1-55.

- Connolly, J.A.D., 1990. Multivariable phase diagrams: an algorithm based on generalized thermodynamics. *Am. J. Sci.* **290**, 666-718.
- Kerrick, D.M., Connolly, J.A.D., 2001. Metamorphic devolatilization of subducted marine sediments and the transport of volatiles into the Earth's mantle. *Nature* **411**, 293-296.
- Connolly, J.A.D., Petrini, K., 2002. An automated strategy for calculation of phase diagram sections and retrieval of rock properties as a function of physical conditions. *J. Metamorph. Geol.* **20**, 697-708.
- Connolly, J.A.D., 2005. Computation of phase equilibria by linear programming: A tool for geodynamic modeling and its application to subduction zone decarbonation. *Earth Planet. Sci. Lett.* **236**, 524-541.

Caddick, M.J., Thompson, A.B., 2008. Quantifying the tectono-metamorphic evolution of pelitic rocks from a wide range of tectonic settings: Mineral compositions in equilibrium. *Contrib. Mineral. Petrol.* **156**, 177-195.

### Perple\_X:

- Minimises free energy of multiple phase configurations to identify the 'most stable' set.
- Can work with complex phases by breaking solution-space into a number of discrete 'phases' (pseudo-compounds).
- Is very flexible, generating:
  - PT projections
  - Pseudosections
  - Compatability diagrams (e.g. AFM triangles)
  - Mixed-variable diagrams
- Is very flexible, allowing:
  - Complex saturation heirarchies
  - Transformation of components
  - Use of numerous EOS for fluids involving H<sub>2</sub>O, CO<sub>2</sub>, NaCl
  - · Easy extraction of sub-systems from large solution model files

#### Perple\_X:

- Permits easy extraction of secondary data:
  - Modes of phases
  - · Compositions of phases
  - Phase or bulk-rock density
  - Other phase or bulk-rock properties (heat capacity, enthalpy, bulk modulus)
  - Estimated secondary properties (P- and S-wave velocities)
- Is a modular collection of many (> 10) sub-programs which can generate a bewildering array of files...
- Requires minimal user input.
- Has the ability to confuse at first.
- · But is actually incredibly simple to use.
- Is much more efficient than it used to be.
- But can still take a long time to calculate a diagram.



















# Solution models for complex phases Phases of variable composition, (solution phases), are invariably described as a mixture of s real or hypothetical endmembers, for which data is tabulated. The problem is then to formulate a solution model that describes the Gibbs energy of such a solution phase in terms of these endmembers. Such models consist of three components $G = G^{mech} + G^{conf} + G^{excess}$ where $G^{mech}$ is the energy arising from mechanically mixing of the endmembers, G<sup>conf</sup> is the energy expected to arise from theoretical entropic considerations, and $G^{excess}$ is a component that accounts for the energetic effects caused by distortions of the atomic structure (e.g., strain) of the chemical mixing process or, in some cases, simply error in G<sup>conf</sup>. (Connolly, 2006)

Three main types of solution model:

- Ideal ( $\Delta S_{\text{config}}$  considered,  $\Delta G_{\text{excess}}$  not considered) Regular ( $\Delta S_{\text{config}} \& \Delta G_{\text{excess}}$  considered)
- Asymmetric (as above, minima not at  $X_i = 0.5$ )

Three levels of complexity:

- Single site mixing (e.g. Mg Fe in olivine; Mg, Fe, Ca, Mn mixing in garnet)
- Multiple site mixing (e.g. Mg, Fe, Ca, Mn & Fe3+ -Al in garnet)
- Multiple site mixing requiring charge balance (e.g. NaSi <sup>3</sup>CaAl in plagioclase feldspar)











So in an ideal solution, *G* is at a minimum at  $X_i = 0.5$  and a stable solid solution is predicted (e.g. forsterite – fayalite). *G* at the minima is strongly controlled by *T*.











### Further reading:

### Derivation of configurational entropy terms:

- Jamie Connolly's course notes (lecture 6)
- Spear, 1993, Metamorphic phase equilibria and *P-T-t* paths, first few pages of chapter 7.

### Regular solution models:

• Powell & Holland, 1993, American Mineralogist, v78, 1174-1180.

### Van Laar solution models:

• Holland & Powell, 2003, Contributions to Mineralogy & Petrology, v145, 492-501.